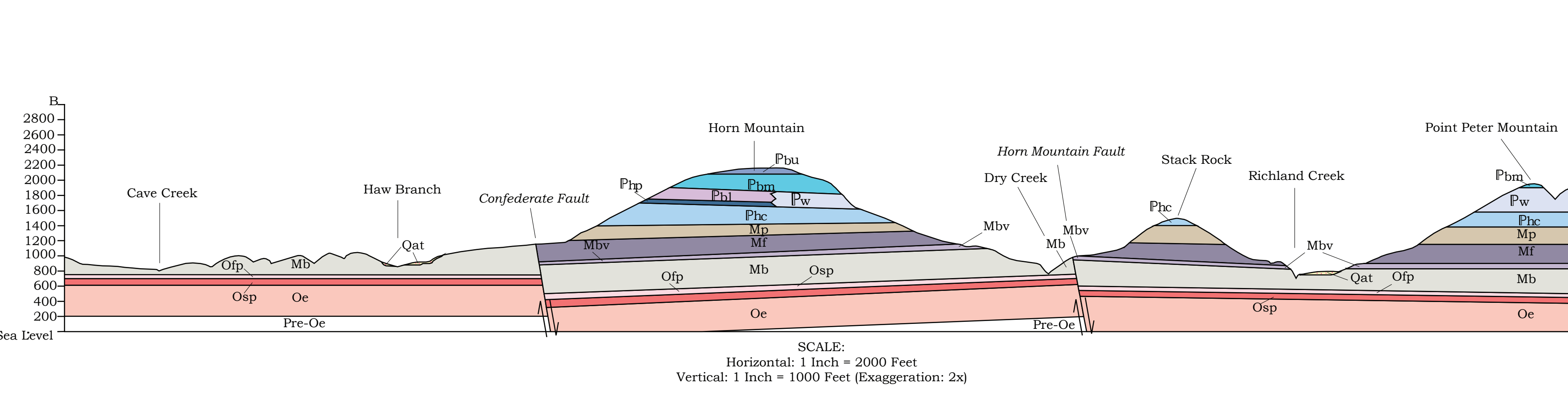
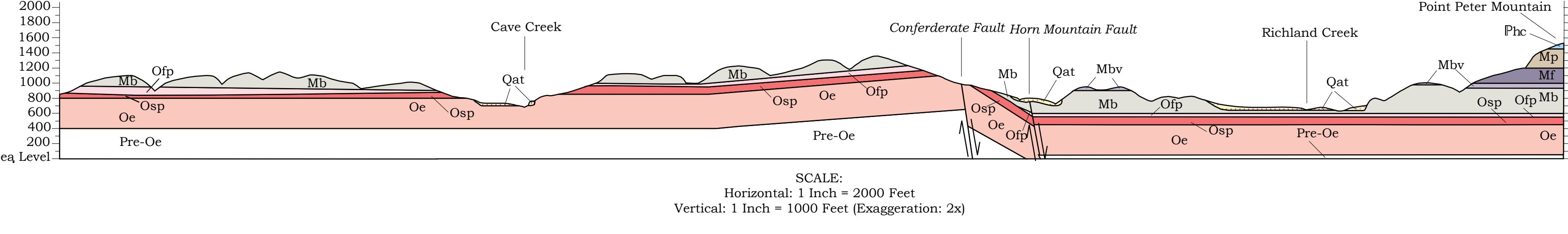
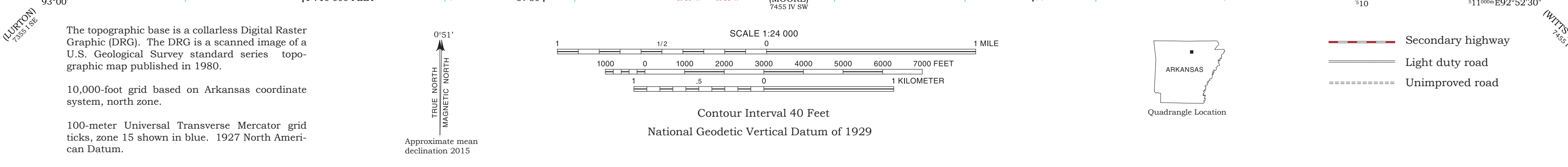
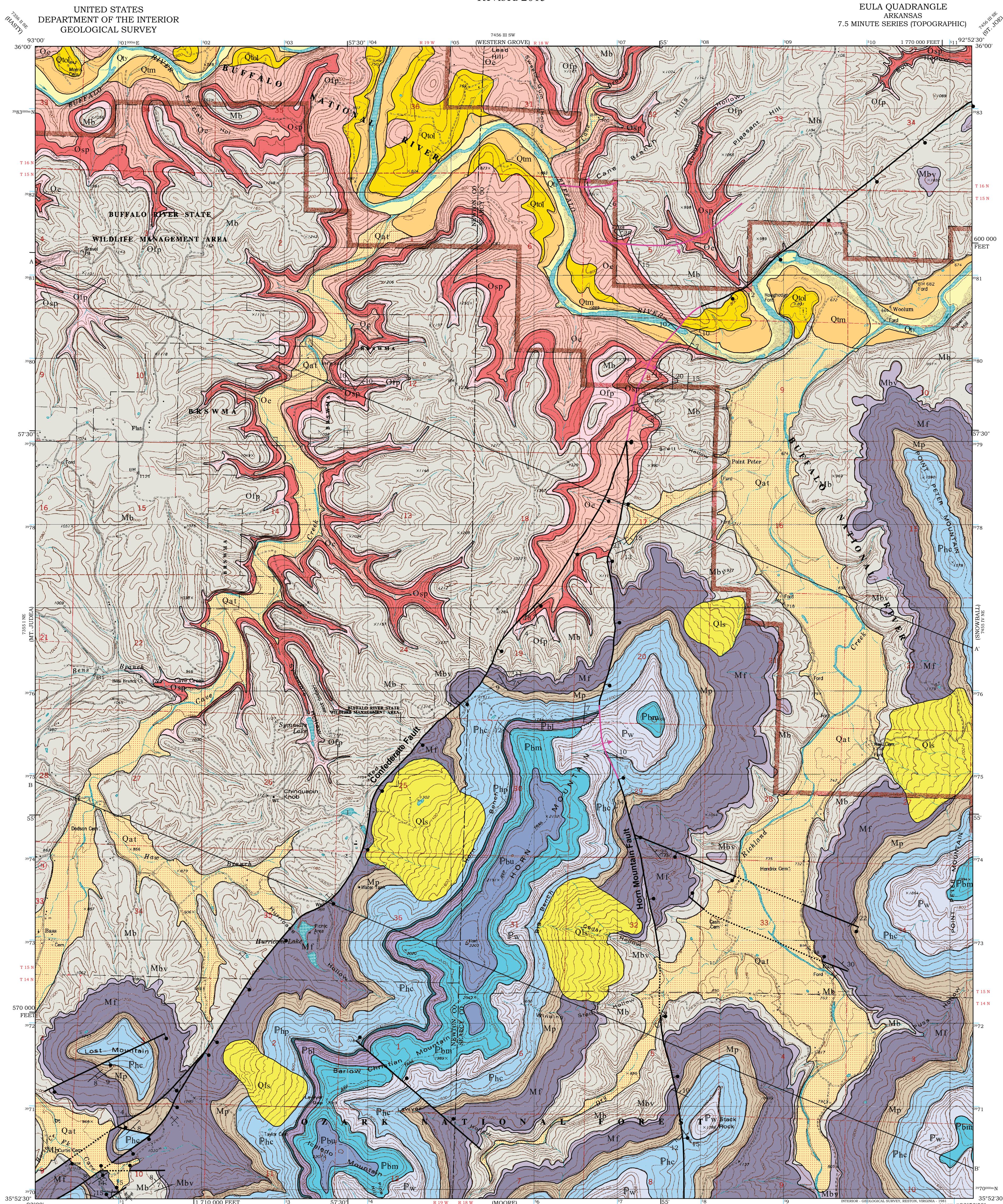
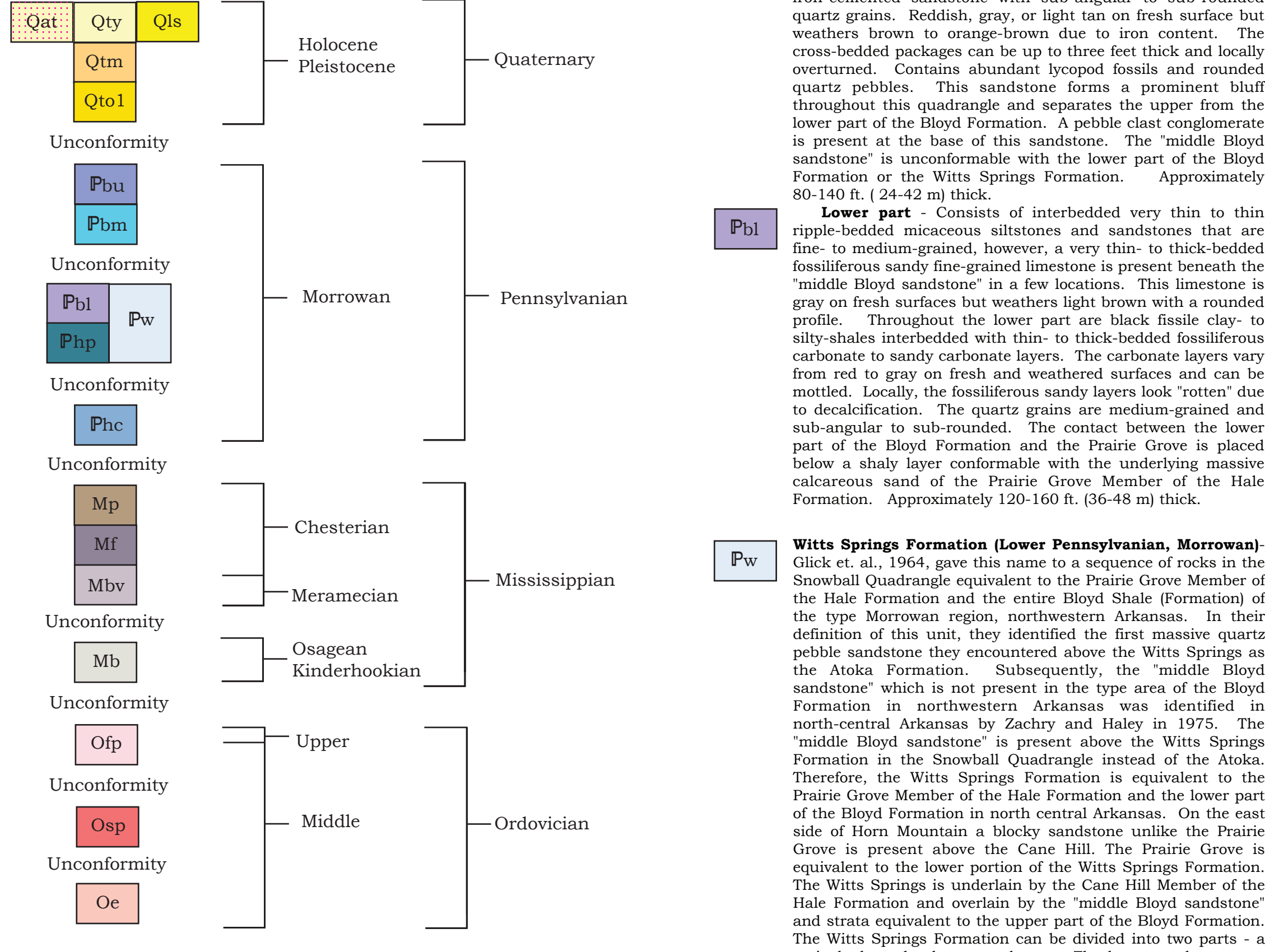


Geologic Map of the Eula Quadrangle, Newton and Searcy Counties, Arkansas

Angela K. Chandler and Scott M. Ausbrooks
2003
Revised 2015



Correlation of Map Units



Introduction

This map illustrates the surface geology of the Eula quadrangle. This quadrangle was previously mapped at a 1:62,500 scale by Glick and Freese in 1965 and Haley in 1976 for the Geologic Map of Arkansas. The geology was mapped at a 1:24,000 scale from 2002-2003 for the National Cooperative Geologic Mapping Program through STATEMAP. During 2014-2015, the terrace deposits along the Buffalo National River were mapped with funding provided by the National Park Service. Approximately eight miles of the Buffalo National River are in this quadrangle and are managed by the National Park Service.

Approximately 2,400 feet (731 meters) of Middle Ordovician to Pennsylvanian age strata crop out in this quadrangle. The Ordovician dolomite, sandstone, and limestone units are present mainly along the Buffalo National River and Cave Creek in the northwestern portion of the map where they are upthrown due to the Confederate Fault. The Boone Formation forms a dissected Springfield Plateau surface over a portion of the map. Upper Mississippian to Pennsylvanian sandstone, shale, and limestone form steep slopes and bluffs on Horn and Point Peter Mountains.

Quaternary alluvium and terrace deposits are present in the valley of the Buffalo National River and its tributaries. Two terrace levels are well developed along the Buffalo River: a younger and a medial. Older terraces are located 80-100 feet (24-49 meters) above the river in most meander bends. This area was prospected in the late 1800's and early 1900's for lead and zinc. Mines and prospect pits are present throughout the quadrangle, however, they are not shown in the National Park since they are considered sensitive park resources.

Description of Map Units

Landslide deposits (Quaternary) - Mostly blocks of sandstone derived from the Morrowan units. Also contains shale slumps from Morrowan and upper Mississippian units.

Alluvium and terrace deposits (Quaternary) - Unconsolidated clay, silt, sand and gravel including deposits on one or more terrace levels.

Young terrace and active channel deposits (Quaternary) - Unconsolidated clay, silt, sand and gravel in gravel bars and sandy point bar deposits along the Buffalo River. Primarily clay, silt, and sand in youngest terrace above the river. The tops of terraces are generally hummocky and tree-covered but can be flat and commonly dissected by tributaries. Approximately 20-30 ft. (6-9 m) thick.

Medial terrace and alluvial deposits (Quaternary) - Unconsolidated clay, silt, sand, gravel, and cobbles located approximately 40 ft. (12 m) above the Buffalo River. The contact with the underlying young terrace is located at a rise that is approximately 15-20 ft. (4-6 m) high and usually coincides with the edge of the riparian zone along the river. Ranges in thickness from 20-60 ft. (6-18 m).

Old terrace and alluvial deposits (Quaternary) - Unconsolidated gravel deposits on ridges above the Buffalo River. Deposit consists of coarse sand to cobble sized sub-angular to rounded chert and sandstone. It is located approximately 80-100 ft. (24-30 m) above the river and ranges up to 160 feet (48 m) above the river. Thickness variable.

Blond Formation (Lower Pennsylvanian, Morrowan) - In this quadrangle, the individual members within the Blond Formation cannot be recognized because the Brentwood and Kessler Limestones are either missing or have become shaly and sandy. There are no other "marker zones" to divide the section into the recognizable members known from the type section in northwest Arkansas. Therefore the Blond Formation is divided informally into lower and upper parts (Hudson et al., 2001) separated by the "middle Blond sandstone" (Zachary and Haley, 1975). Approximately 240-340 ft. (73-103 m) thick.

Upper part - Consists of interbedded thin ripple-bedded to thick-bedded micaceous sandstones and shales. The sandstones consist of fine- to coarse-grained sub-angular to sub-rounded quartz. They are light brown to gray on fresh surfaces but weather dark gray. The shales are dark gray to black on fresh and weathered surfaces. This interval contains many trace fossils, load features, and ball and pillow structures. Approximately 40-140 ft. (12-42 m) thick.

"middle Blond sandstone" - A thin to massive-bedded, medium- to coarse-grained, cross-bedded quartz- or iron-ore-sandstone with sub-angular to sub-rounded quartz grains. Reddish, gray, or light tan on fresh surface but weathers brown to orange-brown to iron content. The cross-bedded packages can be up to three feet thick and locally overturned. Contains abundant lycopod fossils and rounded quartz pebbles. This sandstone forms a prominent bluff throughout this quadrangle and separates the upper from the lower part of the Blond Formation. A pebble-clast conglomerate is present at the base of this sandstone. The "middle Blond sandstone" is unconformable with the lower part of the Blond Formation or the Wits Springs Formation. Approximately 80-140 ft. (24-42 m) thick.

Lower part - Consists of interbedded very thin to thin ripple-bedded micaceous sandstones and sandstones that are fine- to medium-grained, however, a very thin- to thick-bedded fossiliferous sandy fine-grained limestone is present beneath the "middle Blond sandstone" in a few locations. This limestone is gray on fresh surfaces but weathers light brown with a rounded profile. Throughout the lower part is black fissile clay- to silt-shales interbedded with thin- to thick-bedded fossiliferous carbonate to sandy carbonate layers. The carbonate layers vary from red to gray on fresh and weathered surfaces and can be mottled. Locally, the fossiliferous sandy layers look "rotten" due to decalcification. The quartz grains are medium-grained and sub-angular to sub-rounded. The contact between the lower part of the Blond Formation and the Prairie Grove Member of the Hale Formation is approximately 120-160 ft. (36-48 m) thick.

Wits Springs Formation (Lower Pennsylvanian, Morrowan) - Glick et al., 1964, gave this name to a sequence of rocks in the Snowball Quadrangle equivalent to the Prairie Grove Member of the Hale Formation and the entire Blond Shale (Formation) of the type Morrowan region in northwestern Arkansas. In their definition of this unit, they identified the first massive quartz pebble sandstone they encountered above the Wits Springs as the Anoka Formation. Subsequently, the "middle Blond sandstone" which is not present in the type area of the Blond Formation in northwestern Arkansas was identified in north-central Arkansas by Zachary and Haley in 1975. The "middle Blond sandstone" is present above the Wits Springs Formation in the Snowball Quadrangle instead of the Anoka. Therefore, the Wits Springs Formation is equivalent to the Prairie Grove Member of the Hale Formation and the lower part of the Blond Formation in north central Arkansas. On the east side of Horn Mountain a blocky sandstone unlike the Prairie Grove is present above the Cane Hill. The Prairie Grove is equivalent to the lower portion of the Wits Springs Formation. The Wits Springs is underlain by the Cane Hill Member of the Hale Formation and overlain by the "middle Blond sandstone" and strata equivalent to the upper part of the Blond Formation. The Wits Springs Formation can be divided into two parts: a main body and a lower sandstone. The lower sandstone was not mapped separately due to a 40 foot contour interval. Approximately 200-320 ft. (60-97 m) thick.

Main body - Consists mostly of thin- to medium-bedded sandstone and interbedded clay-shale. The sands are very fine to medium grained and usually ripple bedded near the top of the unit. Calcareous fine- to medium-grained fossiliferous sandstones with sub-angular to rounded grains are present and contain clay pebble clasts and fine- to coarse-grained quartz pebbles. Gray on fresh surfaces but weathers brown or dark gray. Approximately 170-290 ft. (51-88 m) thick.

Lower sandstone - A massive coarse-grained iron-ore-sandstone with sub-angular to sub-rounded quartz grains. Sometimes friable. White to yellow on fresh surfaces but weathers a light brown. Contains plant fragments, iron banding, stylolites and root marks. This unit has a blocky appearance and forms a prominent bluff on the east side of Horn Mountain and the southwest side of Point Peter Mountain. A dark-gray shale pebble conglomerate is present at the base of the sandstone. The lower sandstone is unconformable with the Cane Hill Formation. 0- approx. 30 ft. (0-9 m) thick.

Hale Formation (Lower Pennsylvanian, Morrowan) - The Hale Formation consists of two members, the Prairie Grove Member and the Cane Hill Member. Approximately 180-340 ft. (54-103 m) thick.

Prairie Grove Member - A fine- to coarse-grained quartz sandstone with varying amounts of carbonate, crinoid fragments and quartz pebbles. Reddish-gray to brown or mottled on a fresh surfaces but weathers dark-reddish-brown. Bedding varies from thin to massive and exhibits a rounded weathering profile. This unit is a prominent bluff feature that often contains cross-beds and a pitted surface that is referred to as honcomorph weathering. The base of the Prairie Grove Member contains a fossiliferous quartz pebble conglomerate that contains clasts of shale, siltstone, and sandstone as large as almost one foot in diameter. The Prairie Grove Member is unconformable with the Cane Hill Member. 0- approx. 40 ft. (0-12 m) thick.

Cane Hill Member - A gray to black fissile clay- to silt-shale containing iron nodules and small limonitic box work fragments. Varies from black to dark gray on fresh surfaces and light gray and light orange-brown on weathered surfaces. Thin-bedded ripple-bedded siltstones and sandstones are present above the clay shale. Trace fossils are abundant. A 5-20 foot thick sandstone is present beneath the typical Cane Hill block in this quadrangle. It is thin to thick bedded, fine grained, cross bedded, and contains stylolites, plant fragments, and iron banding. One outcrop exposes the contact with the Pitkin as interbedded sandstone with underlying oolitic Pitkin. The Cane Hill Member is unconformable with the Pitkin Limestone. Approximately 180-300 ft. (54-91 m) thick.

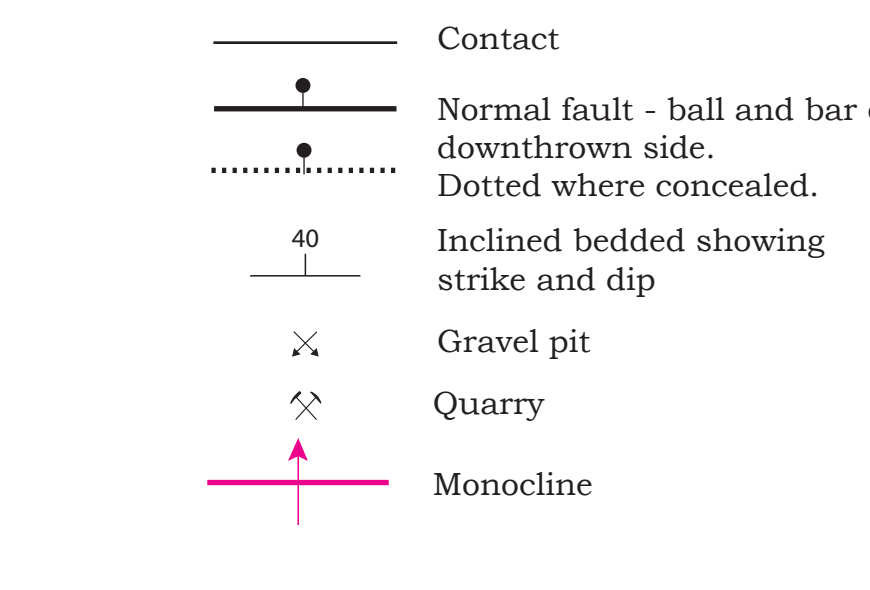
Pitkin Limestone (Upper Mississippian, Chesterian) - A fine- to coarsely-crystalline locally fossiliferous limestone containing crinoid fragments, *Archimedes* bryozoa, gastropods, coral (trigone and colonial) and oolites. Varies from light gray to dark gray on fresh surfaces but usually weathers light or medium gray. Medium to massive-bedded. Often has a petioliferous odor on freshly broken surfaces. A black shale occurs at the top of the Pitkin just beneath the Cane Hill Member of the Hale Formation at a few localities. It is distinctive from the Cane Hill in that it is not silty and does not contain box work fragments. No fossils were found from this shale interval. The Pitkin Limestone is conformable with the Fayetteville Shale. Approximately 120-240 ft. (36-73 m) thick.

Fayetteville Shale (Upper Mississippian, Meramecian) - A black fossil clay shale. Alternating beds of mudstone and shale occur in the upper portion. Black chert is present within the matrix. The matrix beds in the upper portion of this unit form resistant and commonly form deep ledges. Septarian concretions are present near the base of the shale. Very thin fine-grained sandstones are interbedded within the lower portion of the shale. This sandstone dikes with flow structures on both sides are present on the west side of Point Peter Mountain in the lower portion of the shale. The Fayetteville Shale is conformable with the underlying Batesville Sandstone. Approximately 120-280 ft. (36-85 m) thick.

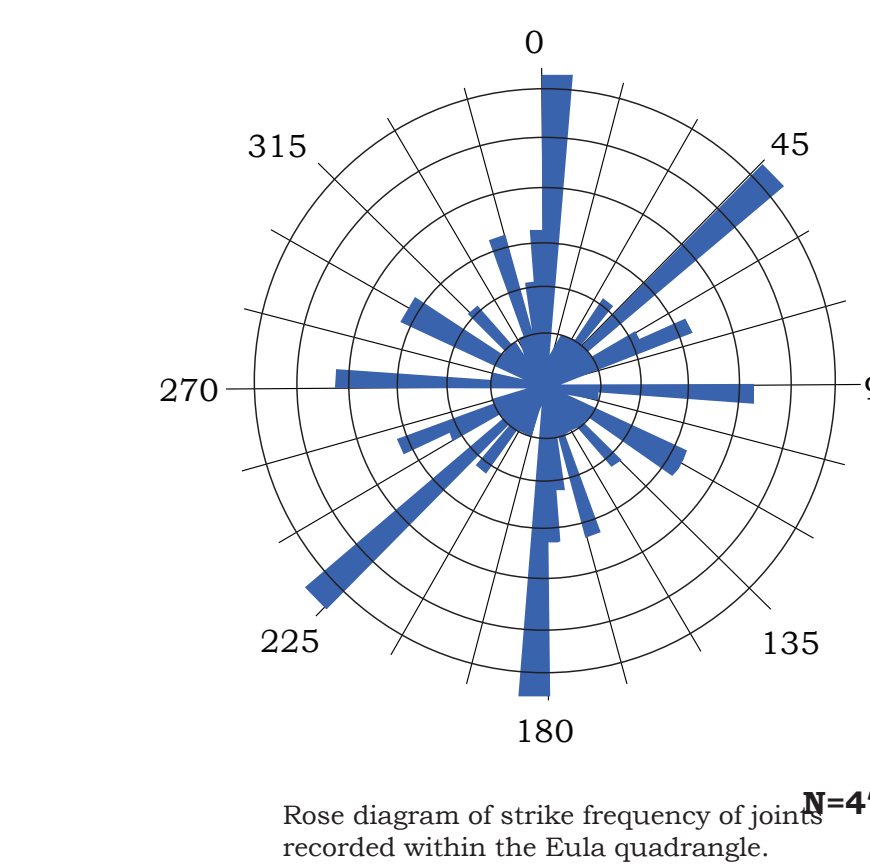
Batesville Formation (Upper Mississippian, Chesterian) - A very fine- to medium-grained, sub-angular, moderately sorted, iron-ore-sandstone. Thin- to medium-bedded. Light brown to cream-colored on fresh surfaces. Weathers light to dark gray. Minor amounts of sandstone are present in this quadrangle. This interval consists mostly of the Hindsville Limestone Member. The Batesville Sandstone is unconformable with the Boone Formation. Approximately 5-40 ft. (1.5-12 m) thick.

Hindsville Limestone Member - A thin-bedded, fine- to coarsely-crystalline limestone. Light to dark gray on fresh surfaces but generally weathers light gray or brown. Typically has a strong petroliferous odor on freshly broken surface. The limestones are fossiliferous and/or oolitic, contain pyrite and are sometimes interbedded with thin layers of clay shale and thin beds of siltstone to fine-grained sandstone. A breccia containing angular chert and limestone fragments is present at the base of this interval at some localities.

Symbols



Joint Frequency



Boone Formation (Lower Mississippian, Osagean and Kinderhookian) - Coarse-grained fossiliferous and fine-grained limestones interbedded with massive and bedded chert. Light to medium gray on fresh surfaces but usually weathers dark gray. The chert varies in color from light gray to dark gray. Springs and sinkholes are abundant. In this area the Boone Formation exhibits an undulating topography that tends to form steep hillsides separated by ravine-like drainages. Approximately 220-400 ft. (67-121 m) thick.

Short Creek Chert - A thin to massive cross-bedded, oolitic crinospirifer and oolitic bryonite in the upper part of the formation. White to gray on fresh and weathered surfaces. Easily recognized by a chalky appearance and in some places a concave weathering profile. Locally, intervals are durable while other intervals are friable. The Short Creek Chert is present at Eula in the section of Boone exposed on the northeast side of Richland Creek at Eula fort. 0- approx. 8 ft. (0-2.4 m) thick.

St. Joe Limestone Member - A medium-grained thin-bedded crinoid limestone containing very thin shaly limestones. Dark gray to reddish in color but sometimes with green mottling on fresh surfaces. Typically weathers a medium to dark gray color. Sometimes contains phosphate nodules near the lower contact. Approximately 0-30 ft. (0-9 m) thick.

Dasil sandstone - A fine- to medium-grained, moderately sorted, sub-rounded to rounded, iron- or quartz-cemented sandstone. White to light gray and tan on fresh surfaces with a salt and pepper appearance that is commonly blanchy due to iron staining. Weathers tan to white. Thin to thick bedded but most often seen as float. Contains phosphate pebbles and angular white and light-gray chert fragments. This unit yields abundant conodonts. The basal sandstone is unconformable with the Lafayette or the Fennelle Limestone. 0- approx. 5 ft. (0-1.5 m) thick.

Lafayette Limestone (Silurian) - A sparsely fossiliferous finely-crystalline limestone. Medium gray with small red blebs on fresh surfaces but weathers light gray. Commonly contains pyrite cubes near the upper contact. Thin to thick bedded with stylolites along bedding planes. This limestone is present on the north hillside along County Road 252 after crossing Cave Creek heading east. It is also present in the head of the drainage in Sec. 19, T15N, R18W on the north edge of Horn Mountain. Unconformable with the Fennelle Limestone. 0- approx. 15 ft. (0-4 m) thick.

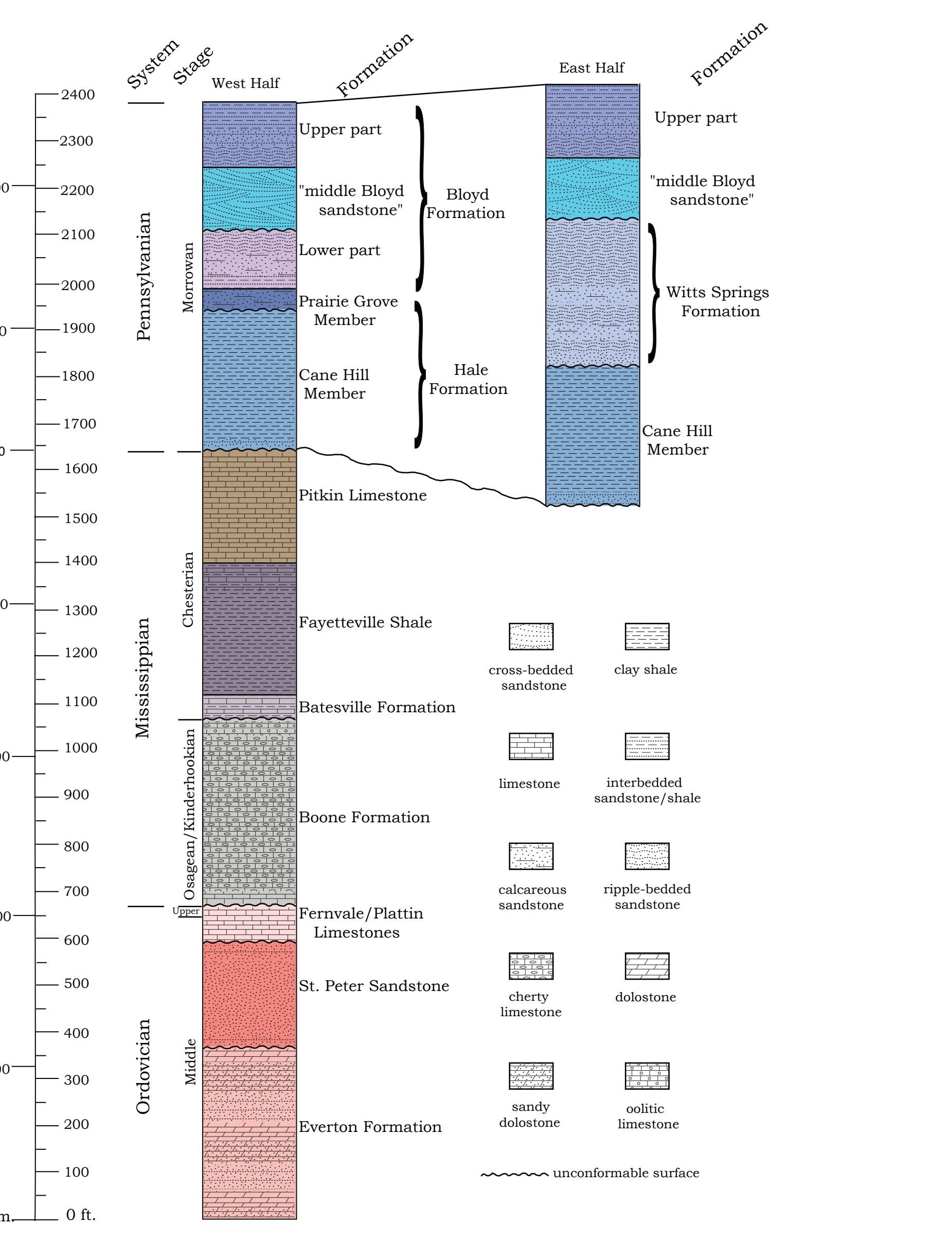
Fennelle Limestone (Upper Ordovician) - A medium- to coarsely-crystalline crinoid limestone. Medium to massive bedded. White to light gray with a pink to reddish tint or mottling on fresh surfaces but weathers a dark-gray. Contains nautiloids, barrel-shaped crinoids, and brachiopods that are accentuated on a weathered surface. Commonly contains pyrite. Locally cross-bedded when beds are massive. On a weathered slope the Fennelle Limestone occurs as rounded masses that are usually friable. The Fennelle is unconformable with the Plattin Limestone. 0- approx. 15 ft. (0-4 m) thick.

Plattin Limestone (Middle Ordovician) - A thin- to thick-bedded micritic limestone that locally displays a sugary texture. Light gray to dark gray on fresh surfaces and weathers white to dark gray. A dolomitic interval is present at the top of the formation. The Plattin Limestone is unconformable with the St. Peter Sandstone or Everton Formation in this quadrangle. 5- approx. 80 ft. (15-24 m) thick.

St. Peter Sandstone (Middle Ordovician) - A thin- to thick-bedded very fine- to fine-grained sandstone. White to green on fresh surfaces but weathers light gray-green to brown. Green shale clasts are present and locally weather to give the sandstone a green color. Contains a calcareous matrix that weathers a light brown. The quartz grains are rounded. Green siltstones and shales are interbedded with the sandstone. Contains vertical trace fossils referred to as Scolithus by Adams et al., 1904, that weather to resemble scissels in cross-section view. The St. Peter Sandstone is unconformable with the Everton Formation. Approximately 20-160 ft. (6-48 m) thick.

Everton Formation (Middle Ordovician) - Very fine- to fine-grained sandy and clay dolostones that are thin- to massive-bedded. Thin to medium beds of fine- to medium-grained quartz sandstone are common and similar to the overlying St. Peter Sandstone. Medium to dark gray on fresh surfaces but usually weathers light gray. Approximately 10-360 ft. (3-109 m) exposed at the surface.

Stratigraphic Column



References

Adams, G.I., Purdie, A.H., Burchard, E.F., and Ulrich, E.O., 1904. Zinc and lead deposits of northern Arkansas with a section on the determination and correlation of formations. U.S. Geological Survey Professional Paper, no. 24, 118 p.

Glick, E.E., Freese, S.E., and Gordon, M.J., 1964. Wits Springs Formation of Morrow age in the Snowball quadrangle, north-central Arkansas: in Contributions to stratigraphy, U.S. Geological Survey Bulletin 1194-D, p. D1-D16.

Glick, E.E. and Freese, S.E., 1965. Geologic map of the Snowball quadrangle, Newton and Searcy Counties, Arkansas: U.S. Geological Survey, Geologic quadrangle maps of the United States, Map GQ-425, 1 sheet, 3 p.

Haley, B.R., 1976. Geologic map of the Snowball quadrangle, 15-minute series: Arkansas Geological Survey Geologic Worksheet, 1 sheet.

Hudson, M.R., Murray, K.E., and Pezzutti, D., 2001. Geology of the Jasper quadrangle, Newton and Boone Counties, Arkansas: U.S. Geological Survey Miscellaneous Field Studies Map MF-2356.

McKnight, E.J., 1935. Zinc and lead deposits of northern Arkansas. U.S. Geological Survey Bulletin 853, 31 p.

Zachary, D.L. and Haley, B.R., 1975. Stratigraphic relationships between the Blond and Anoka Formations (Pennsylvanian) of northern Arkansas: in Contributions to geology of the Arkansas Ozarks, Arkansas Geological Commission, Miscellaneous Publication 12, p. 96-106.

Acknowledgments: This map was produced for STATEMAP, Cooperative Agreement Award O2HQAG0017, a matching funds grants program with the US Geological Survey under The National Cooperative Geologic Mapping Program. Thanks to Corbin Cannon and Danny Ratins for their assistance with mapping the terrace deposits along the Buffalo National River. Special thanks to the National Park Service, U.S. Forest Service, Game and Fish Commission, and to private landowners who graciously allowed access to their property.

Limitations: This map, like all geologic maps, is based on interpretations which were made from the data available at the time it was created. As work continues and new data is collected, the contacts and structures depicted on this map may be changed.

For the latest edition of this and other AOS maps and publications, please call Publication Sales at 801-296-1877, or 800-854-7444, or the Vandell Parham Geology Center, 3815 West Roosevelt Road, Little Rock, Arkansas 72204. This map is available at: http://www.geology.ar.gov/geologicmaps/dgm_24k.htm.

Suggested citation: Chandler, A.K. and Ausbrooks, S.M., revised 2015. Geologic map of the Eula quadrangle, Newton and Searcy Counties, Arkansas: Arkansas Geological Survey Digital Geologic Map, DGM-AR-00269, 1:24,000.